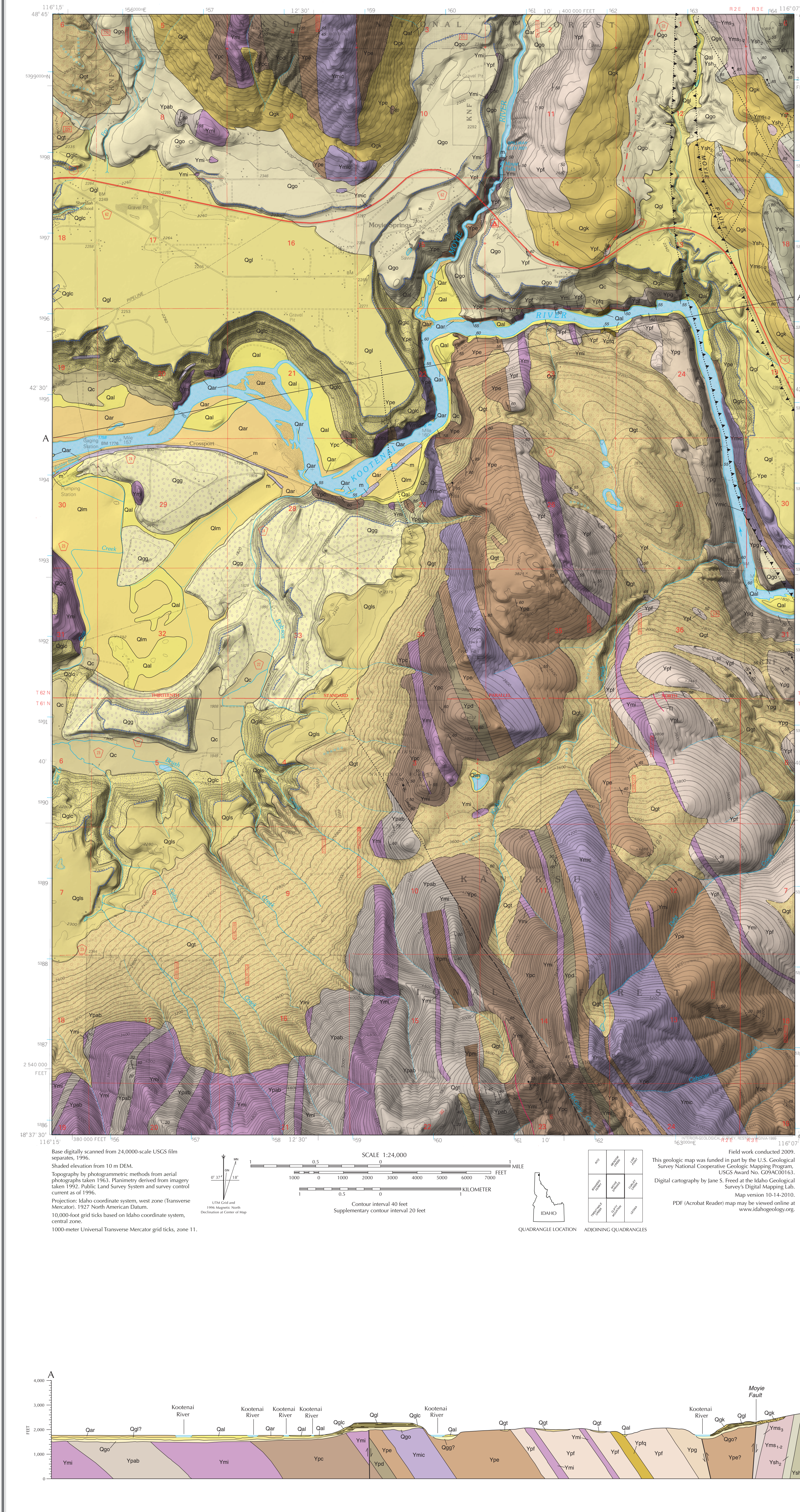


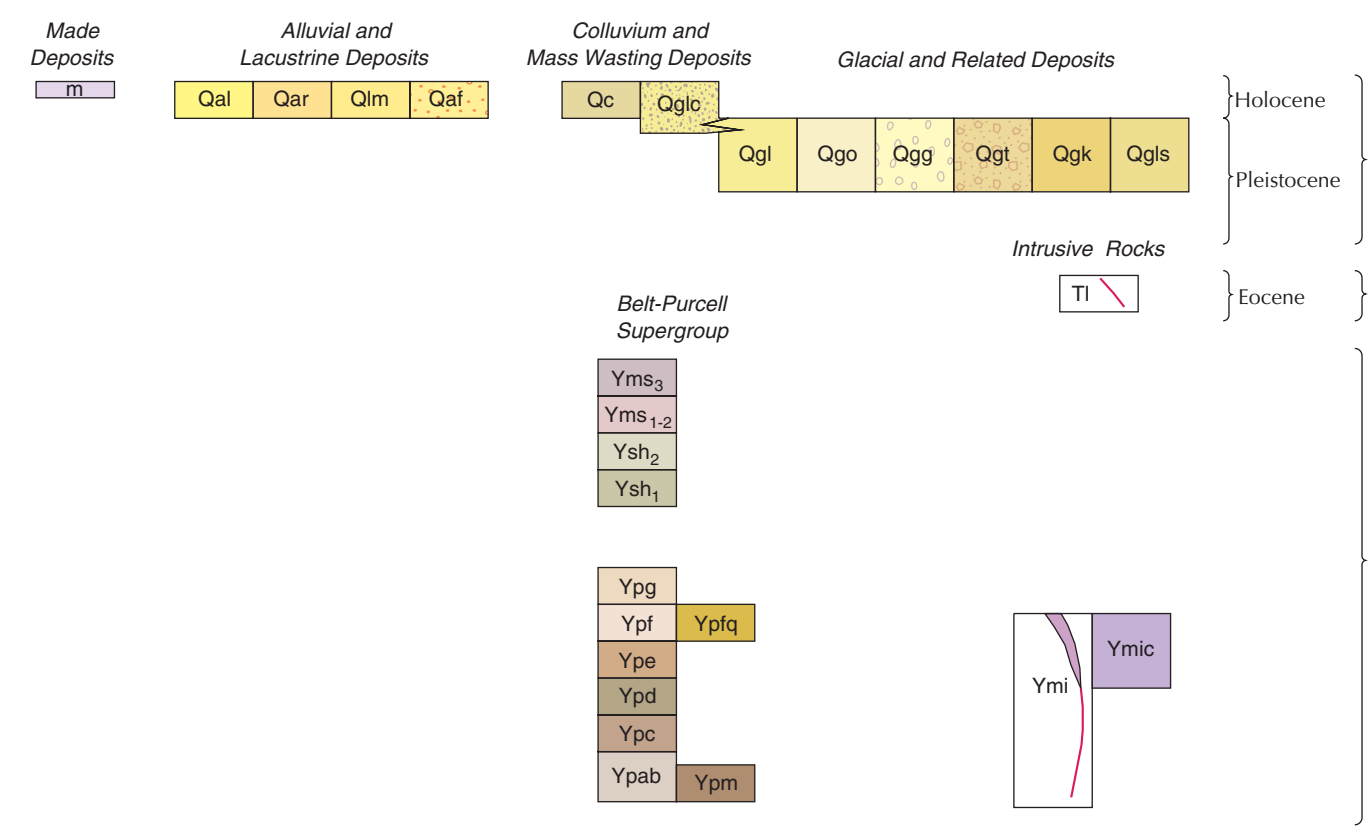
GEOLOGIC MAP OF THE MOYIE SPRINGS QUADRANGLE, BOUNDARY COUNTY, IDAHO

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CORRELATION OF MAP UNITS



INTRODUCTION

Quaternary deposits on this 1:24,000-scale quadrangle were mapped in 1992 and 2008-2009 by R.M. Breckenridge. Bedrock was mapped in 2009 by M.D. McFadden, R.F. Burmester and R.S. Lewis to modify previous mapping (Burmester, 1986) for consistency with unit definitions and contact placements used in more current mapping to the west.

The most abundant rocks in the Moyie Springs quadrangle are low metamorphic grade metasedimentary rocks of the Belt-Purcell Supergroup. Precambrian in age. The oldest rocks underlie the western part of the map, have tops to the east, and host mafic and differentiated sills. And of the youngest strata, on the west flank of the Sylvan anticline east of the Moyie fault (Fig. 1), have tops to the west. Sediments in the lowlands and some at higher elevations date from Pleistocene glaciation and damming of the Kootenai River.

DESCRIPTION OF MAP UNITS

Intrusive rocks are classified according to IUGS nomenclature using normal-zoned values of modal quartz (Q), alkali feldspar (A), and plagioclase (P) on a ternary diagram (Streckeisen, 1976). Mineral modalities are listed in order of increasing abundance for both igneous and metamorphic rocks. Grain size classification of unconsolidated and consolidated sediment is based on the Wentworth scale (Lane, 1947). Bedding thicknesses and laminar type are after McKee and West (1963) and Winston (1986). Thicknesses and distances are given in abbreviation of metric units (e.g., dm=decimeter); some linear dimensions in parentheses are in English units (meters). Within a rock unit description are listed in order of decreasing abundance: Quaternary units are after Chugg and Folsberg (1980) and Weisel (2005).

MAN-MADE DEPOSITS

made land (historical)—Highway fills, railroad rights of way, levees of the Kootenai River and the adjacent alluvial flood plain. Numerous small fills are unmapped.

ALLUVIAL AND LACUSTRINE DEPOSITS

Active river wash (Holocene)—Silt, clay, and sand deposits in the active channel and flood plain confined by levees of the Kootenai River. Most channel substrates represent a modern deposit related to the 1972 completion of Libby Dam upstream in Montana.

Alluvium (Holocene)—Alluvial deposits of the Kootenai and Moyie rivers and tributaries. Moderately sorted to well-sorted silt, sand, and local pebbles and cobbles. Mostly revealed glacial deposits in the river valley and post-glacial colluvium in the surrounding mountains. Schoonover-Ritz-Farnham soils are associated with these very deep, fine clay loams, silt loams, and mucky silt loams in basins and swales and on low terraces, flood plains, and natural levees. Thickness is several to more than 10 m (6 to 30 ft).

Alluvial fan deposits (Holocene)—Mixed pebbles to cobble gravel deposited as fans at the mouths of local drainages. Mostly subangular to angular clasts derived locally from colluvium and glacial deposits on steep slopes. Thickness is several to more than 5 m (6 to 15 ft).

Lacustrine and mud deposits (Holocene)—Organic mud, mud, and peat bogs in poorly drained paleogeological outwash channels and kettle depressions. Interbedded with thin layers of fine sand, silt, and clay. Soils of the Pywell series. Thickness from 1-5 m (3-16 ft).

COLLUVIAL AND MASS WASTING DEPOSITS

Colluvial deposits (Holocene)—Silt, sand, and gravel colluvium. Forms debris fans and colluvial aprons along steep escarpments and gullies of terraces and benches. Includes small unmappable mass movements. This unit mostly in escarpments of Qgl; where mappable. Varied thickness as much as several meters (15 ft).

Glaciolacustrine deposits (Pleistocene to Holocene)—Mixed deposits of silt, sand, and gravel colluvium, slope wash, and small landslides. Steep slopes of reworked and locally transported Qgl. Soils are silt loams of the Weibstone-Crash association. Thickness as much as 10 m (30 ft).

GLACIAL AND FLOOD-RELATED DEPOSITS

Glaciolacustrine deposits (Pleistocene to Holocene)—Massive to well-bedded and finely laminated clay, silt, and sand deposited in Glacial Lake Kootenai at the northern retreating ice margin. Ice deposits are mostly clay, silt, and sand with developed rhythmites and beds of sand and silt and scattered dropstones. Contorted bedding and loading structures are common. This unit forms several prominent terrace levels (200 to an elevation about 2,200 to 2,300 feet) also forms discontinuous terraces in tributary valleys. Mostly well sorted and finely laminated. Overlain by glaciolacustrine outwash deposits on terraces and in tributary valleys. Soils are silt loam and silty sandy loams of the Weibstone-Crash association. Exposed thickness as much as 150 m (500 ft).

Outwash deposits, undivided (Pleistocene)—Silty and sandy gravels. Moderately sorted and rounded pebbles and cobbles. Outwash deposits in this quadrangle are from glaciation of the Moyie River valley that were graded to Glacial Lake Kootenai. Forms a series of terraces at the mouth of the Moyie River. The presence of Glacier Peak tephra (11,200 R.P.) just downstream near Bonner Ferry gives a minimum date for retreat of the ice lobe from the Purcell Trench (Richmond, 1986). Soils are sandy loams and loamy sands of the Sella-Elmira Association. Thickness as much as 20 m (60 ft).

Glacial gravel deposits (Pleistocene)—Coarse cobble and boulder gravel deposits downstream from the junction of the Moyie and Kootenai rivers as they exit from the mountains. Forms a large gravel delta interbedded with glaciolacustrine deposits of Glacial Lake Kootenai. Soils of the Bonner series. Thickness 5-30 m (15-100 ft).

Till deposits (Pleistocene)—Dense silt, pebble, and cobble till with local boulders deposited by the Purcell Trench lobe of the Cordilleran ice sheet. Poorly stratified compact basal till includes ground moraine and some interbedded proglacial deposits. Deposit includes kame terraces and some outwash. Soils include silt loams and gravelly silt loams of the Bend Oreille rock outcrop and the Stien-Pend Oreille associations. Thickness from 1-50 m (3-150 ft) in subsurface.

Glaciolacustrine and ice stagnation deposits (Pleistocene)—Fine silt and sand with interbedded bar and channel sand and gravel deposits. This laminae is thickly bedded sands with some cross bedding. Dunes (giant current ripples) and closed depressions (kettles) are the dominant surface features. Soils are very deep fine sandy loams and loamy sands of the Sella-Elmira association. Thickness 2-20 m (6 to 60 ft).

Kame deposits (Pleistocene)—Poorly stratified and compact silty to sandy boulder lodgement till; locally includes ground moraine and some interbedded proglacial and ice contact and outwash deposits. This unit probably underlies Qgl and records the latest recession of the Purcell Trench ice lobe. Soils include silt loams and gravelly silt loams of the Stien-Pend Oreille association. Thickness varies; may exceed 50 m (160 ft).

INTRUSIVE ROCKS

Lamprophyre dikes (Eocene)—Biotite lamprophyre dikes with 1-2 mm biotite phenocrysts in a fine-grained groundmass; grain size varies between and within dikes. Known only in center east half of map area. Age presumed similar to a lamprophyre dike that cuts the Sandpoint conglomerate in the Elmira quadrangle (Lewis and others, 2007) to the south, which has an ⁴⁰Ar/³⁹Ar age of 47.15 ± 0.24 Ma on biotite (Doughy and Price, 2000).

Mafic intrusive rocks, undivided (Mesoproterozoic)—Fine- to medium-grained, rare coarse-grained hornblende gabbro. Variants range from high color index (CI) hornblende to low CI differentiates that are most common at or near sills as medium-grained quartz diorite or biotite granophyre, but also occur separately. Plagioclase rarely as cm-long laths, quartz as 5 mm blue grains or in cm-diameter granophyre concentrations. Variations are common between intrusions and within single bodies both along and across strata. All within the map area are Moyie sills described by Bishop (1976); they are likely from Holocene magmatic or differentiates intruded into the lower and middle parts of the Purcell Formation (Hoy, 1989). Intrusions are generally concordant but some laterally pinch and swell, include rays of country rock, branch into multiple sills, end abruptly, or gradually pinch out (Bishop, 1976). Sills lower and higher in the section are similar in chemistry (Anderson and Goodfellow, 2000), but they seem to have intruded in at least two separate events. Early intrusions were at shallow levels closely following or were invasive during sedimentation of the Lower Aldridge equivalent Ypab (Hoy and others, 2000; Gorton and others, 2000; Cressman, 1969; Sears and others, 1998; Page and others, 2000). Generation of Ypab was probably synchronous with shallow intrusive events. Higher intrusions stop chilled margins and contact aureoles, evidence they invaded consolidated rock at a later time. Age of sills in Ypab is probably close to U-Pb dates on zircons from near Kimberley, British Columbia about 100 km to the north (1.468 Ga; Anderson and Davis, 1995) and near Plains, Montana, about 190 km south-east (1.47 Ga; Sears and others, 1998). The younger age from the differentiated top Ypab in Ypab of 1.433 Ga (Zartman and others, 1982) has been attributed to metamorphism (Anderson and Davis, 1995).

Mafic intrusive rocks, Crompton C sill (Mesoproterozoic)—Fine- to coarse-grained hornblende gabbro, hornblende, quartz gabbro to quartz diorite. Fine grained variety typically near top and bottom, but also occurs interlayered with coarser phases parallel to upper and lower conformable contacts within the body. Granophyre blotches to 2 cm common near top of unit across most of the map. Hornblende locally about that has acicular amphibole 2 cm long in contact. In common stubby form. Where observed, lower contact chilled. Thickness near center of map about 360 m (1,200 ft) (Bishop, 1976).

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STRUCTURE

The major structure in the map area, although not exposed, is the Moyie fault. It is characterized regionally as an east-vergent thrust, and does juxtapose older rocks on the west against younger ones on the east, consistent with this interpretation. However, those rocks face each other giving the impression that the fault occupies a suture. Support for a synclinal fold geometry comes from existence of a slightly east-verging southward-plunging syncline west of the Moyie fault south of the map area (Miller and Burmester, 2004), an open, northward-plunging syncline east of the fault north of the international border (Brown and others, 1995), and a siver of east-facing upper Belt strata east of the fault near the northeast corner of this map. Latest motion on the Moyie fault was probably down to the west, as was motion on a fault to its west and parallel to it that repeats part of the Purcell Formation and Ypab at the east edge of the map.

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SYMBOLS

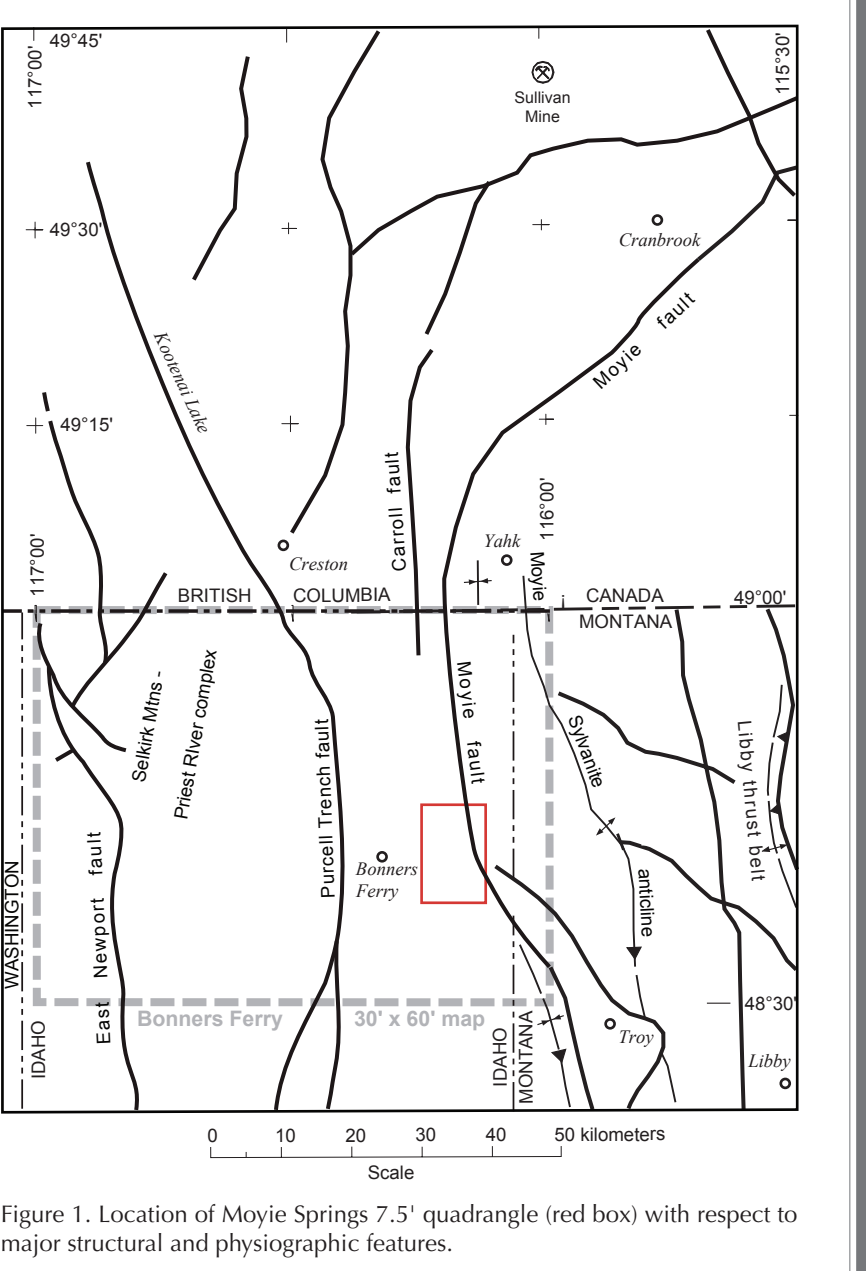
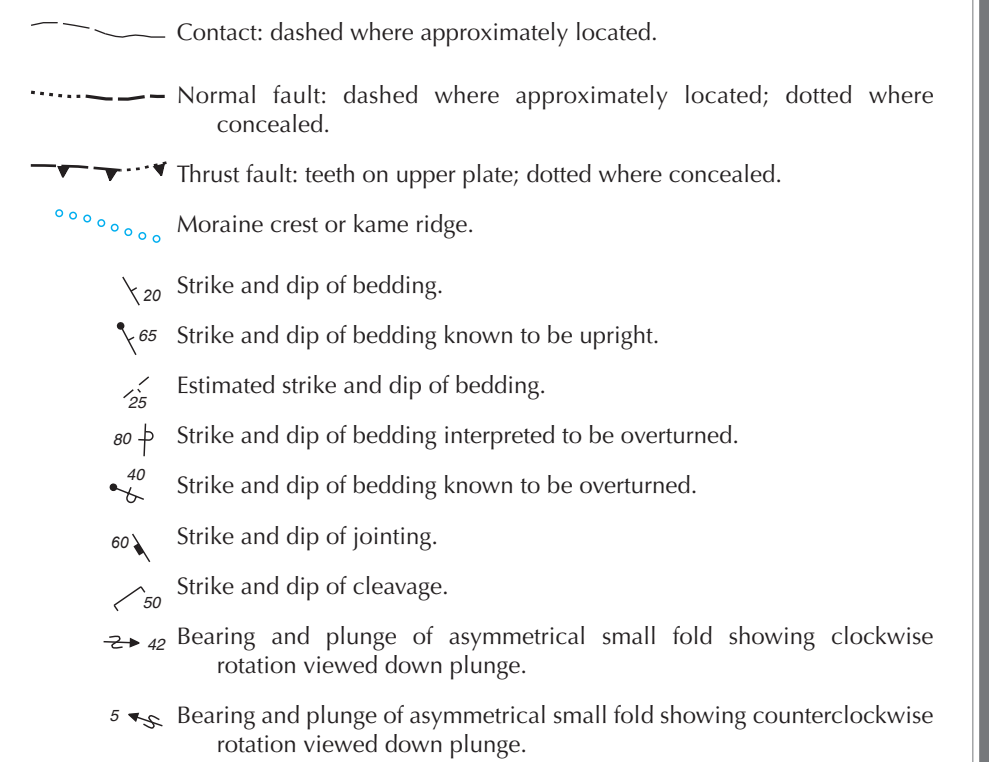


Figure 1. Location of Moyie Springs 7.5' quadrangle (red box) with respect to major structural and physiographic features.